Dynamic Modeling of the 2004 M_w 6.0 Parkfield Earthquake

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Distribution of Peak Acceleration

8 strong ground motion stations < 1 km to the fault

40 station between 1 and 10 km

13 continuous GPS stations

Shakal et al, 2004

Velocity Structure



"Softer" Franciscan assemblage rocks (NE)

"Stiffer" Salinian granitic rocks (SW)

Dynamic Rupture on a Bi-Material Interface

asymmetry in normal stress variations



e.g., Harris and Day (1997, 2005), Rubin and Ampuero (2007) ...

3D Simulation









4.0 sec

Harris and Day, 2005

Laboratory Confirmation

(B)





Xia et al., 2005

Supershear

Aftershock Asymmetry of Micro-Earthquakes



The asymmetry in normal stress variation on a bimaterial interface can relate to the observed aftershock asymmetry on the SAF (Rubin and Ampuero, 2007)

1D Velocity Structure



From Cuscodio et al. (2005)

1D velocity model interpolated from the 3D velocity model of Thurber et al. (2004)

Slip and Static Stress Drop



solve a static problem



Slip Weakening Friction



Andrews, 1976

Stress Drop and Strength Excess





Model B



A nearly constant S ratio was used to determine the strength excess.

Seismicity 1984-2005



Blue: seismicity before the 2004 mainshock

Red: the 2004 mainshock and its aftershocks

Thurber et al., 2006

FEM Calculation Parameters

Fault Strike: 140°

Fault Dip: 90°

Fault Length: 40 km

Fault Width: 15 km (the fault is 500 m below the surface)

Element size: 100 m

Time Step: 0.012 s

Critical Slip-Weakening Distance (D_c): 0.15 m

Initial normal stress on fault: 60 MPa

Dynamic frictional coefficient: 0.3

Initial and static friction coefficients are determined by stress drop and strength excess.

Evolution of Slip Rate, Shear and Normal Stress Changes



Space-Time Plot of Slip Velocity at the Hypocentral Depth



Peak Surface Velocity





Synthetics vs. Data: Cholame





Synthetics vs. Data: Stone Canyon



Synthetics vs. Data: Vineyard Canyon



Fault-Normal Velocities Very Close to the Fault



Misfit for 43 Stations

misfit =
$$\sum_{i=1}^{nstn} \sum_{n=1}^{npts} [v_i^{syn}(t_n) - v_i^{data}(t_n)]^2$$

Custódio et al. (2005)fault-parallel velocity: $2.92 \text{ m}^2/\text{s}^2$ fault-normal velocity: $4.67 \text{ m}^2/\text{s}^2$ vertical velocity: $1.14 \text{ m}^2/\text{s}^2$

Model A

fault-parallel velocity:3.94 m²/s²fault-normal velocity:5.82 m²/s²vertical velocity:1.13 m²/s²

Model B

fault-parallel velocity: $3.30 \text{ m}^2/\text{s}^2$ fault-normal velocity: $5.66 \text{ m}^2/\text{s}^2$ vertical velocity: $0.91 \text{ m}^2/\text{s}^2$

Coseismic Stress Change vs. Seismicity Before (blue) and After (red) the Mainshock



Predicted Coseismic Offset vs. GPS

Model A

Model B



GPS absolute offsets were obtained by taking the difference of averaged positions over 40 - 100 s after the mainshock and 100 s before the mainshock.

Slip Distributions







Conclusions

 Buried slips on the SAF between 5 km and 10 km depth can largely explain both ground motion and GPS observations during the 2004 Parkfield earthquake.

 We found the stress drop to be ~ 10 MPa in the hypocentral region and 1~2 MPa on other slipped regions of the fault.

 The main rupture front propagates bilaterially at almost a constant subshear rupture velocity ~3 km/s.

 Material contrasts across the fault caused significant normal stress variations on the fault, which affects strongly rupture propagation and ground motion.

The large ground motion in Cholame can be attributed to the large dynamic stress drop in the positive (southeast) direction.