

***Heterogeneous Initial Stress Conditions
for Spontaneous Rupture Modeling***

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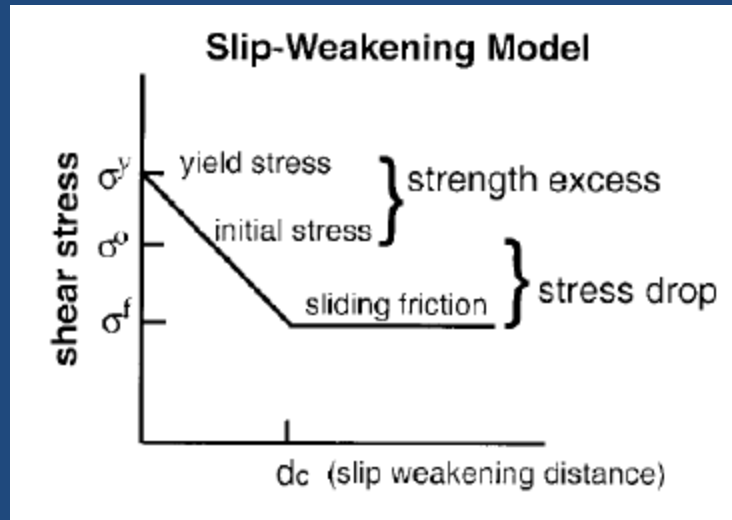
SCEC 100-runs workshop at USC

May 21, 2010

Outline

- **Spatial heterogeneity with 1-point and 2-point statistics**
 - Do we pay enough attention to 1-point statistics?
 - How to constrain them from data?
- **Two-step approach**
 - Quasi-dynamic multi-cycle simulation with RS friction law
 - Full-dynamic single-event simulation with SW friction law

Spontaneous Rupture Modeling *with Slip-weakening Friction Law*



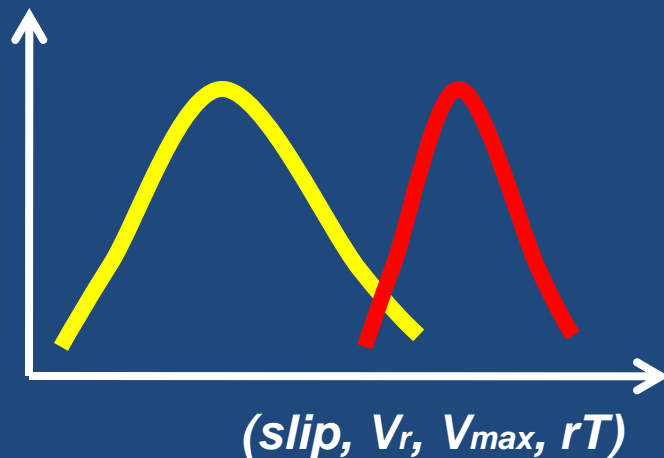
(Ida, 1972; Andrews, 1976)

- Stress drop: from given slip models, or assumed stochastic model (e.g., $k^{-\nu}$)
- Fracture energy: somewhat arbitrary, i.e., S parameter, constant yield stress, strength excess, d_c , etc.

Earthquake Source Statistics

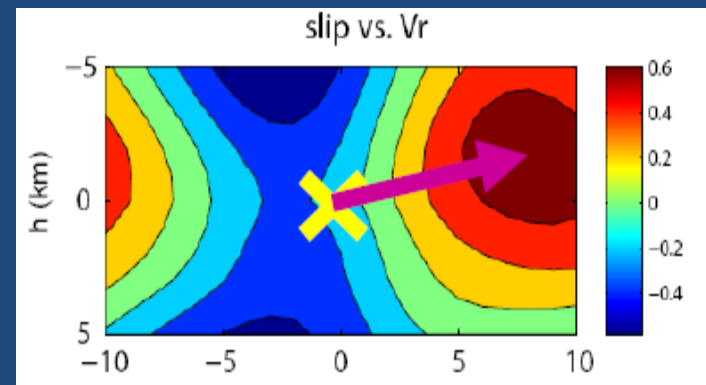
1-Point Statistics

- Scaling of mean slip and sigma with earthquake size
- Supershear and subshear
- Crack-like and pulse-like rupture
- Stick-slip and creeping



2-Point statistics

- *Auto-coherence*: define heterogeneity of source parameters
- *Cross-coherence*: control coupling between different parameters



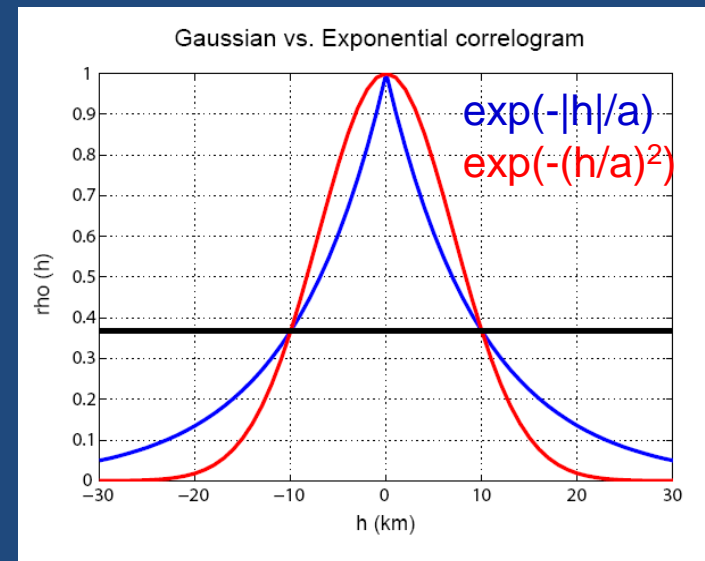
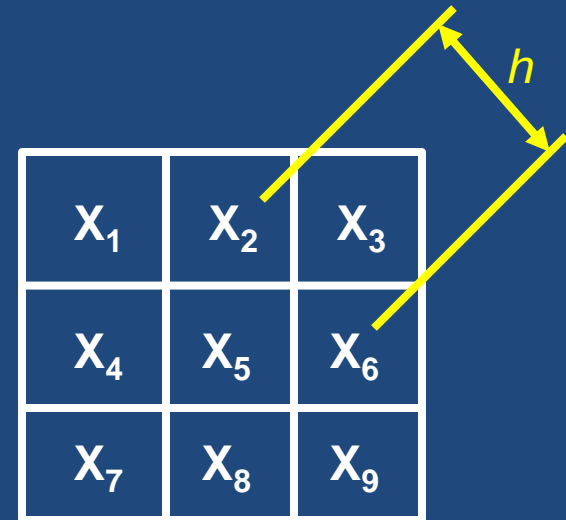
Spatial Random Field Model

$$\mathbf{X} \sim \mathcal{N}(\boldsymbol{\mu}, \boldsymbol{\Sigma})$$

$$\begin{aligned} \boldsymbol{\mu} &= (\mu_1, \mu_2, \mu_3, \dots, \mu_9)^T \\ &= (\mu, \mu, \mu, \dots, \mu)^T \end{aligned}$$

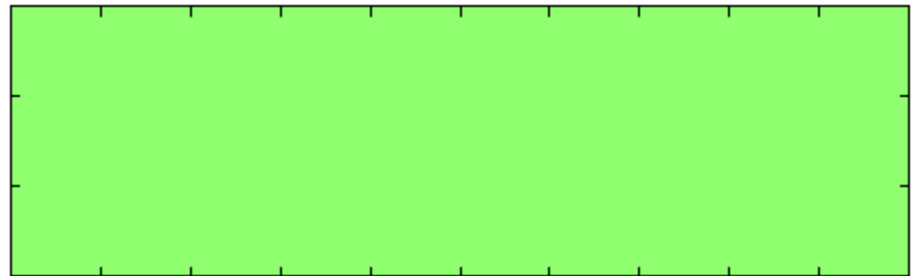
$$\boldsymbol{\Sigma} = \begin{bmatrix} \sigma_1\sigma_1\rho_{11}, & \sigma_1\sigma_2\rho_{12}, & \dots, & \sigma_1\sigma_9\rho_{19} \\ \sigma_2\sigma_1\rho_{21}, & \sigma_2\sigma_2\rho_{22}, & \dots, & \sigma_2\sigma_9\rho_{29} \\ & & \dots & \\ \sigma_9\sigma_1\rho_{91}, & \sigma_9\sigma_2\rho_{92}, & \dots, & \sigma_9\sigma_9\rho_{99} \end{bmatrix}$$

$$= \sigma^2 \begin{bmatrix} 1, \rho_{12}, \dots, \rho_{19} \\ \rho_{21}, 1, \dots, \rho_{29} \\ \dots \\ \rho_{91}, \rho_{92}, \dots, 1 \end{bmatrix}$$

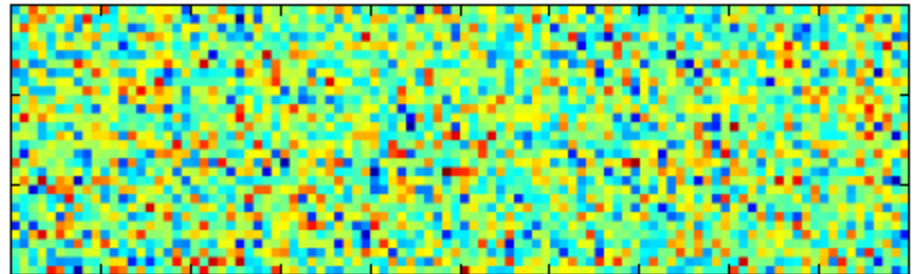


Reproduction of *1-point and 2-point statistics*

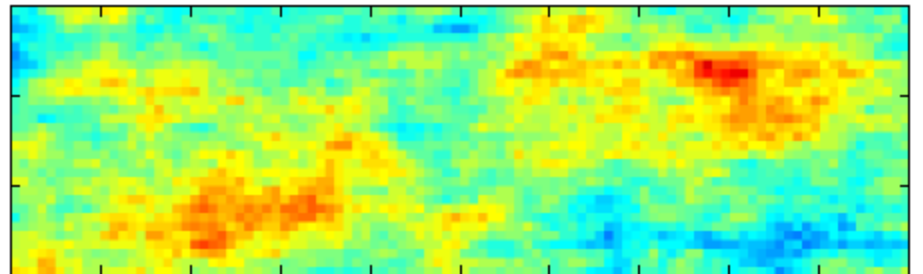
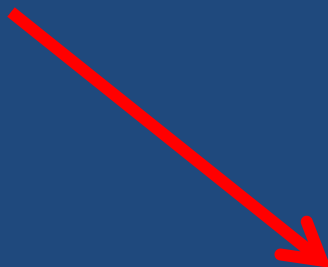
1) $\mathbf{X} \sim \mathcal{N}(\boldsymbol{\mu}, \mathbf{0})$



2) $\mathbf{X} \sim \mathcal{N}(\boldsymbol{\mu}, \sigma^2 \mathbf{I})$

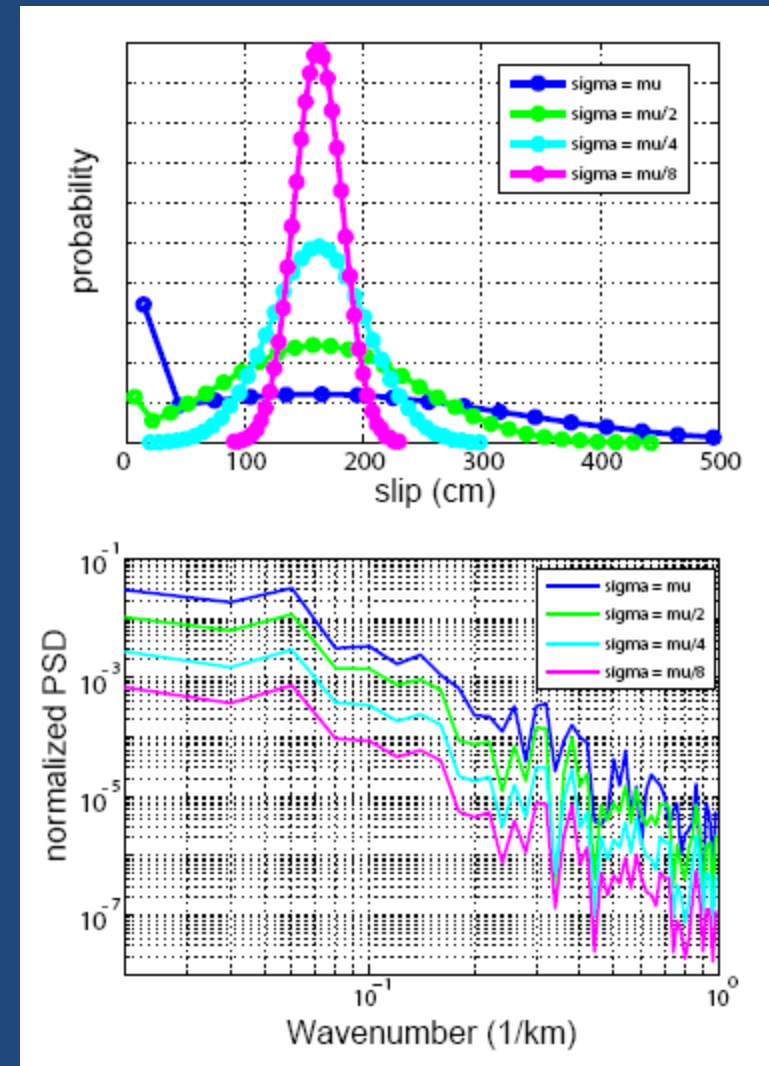
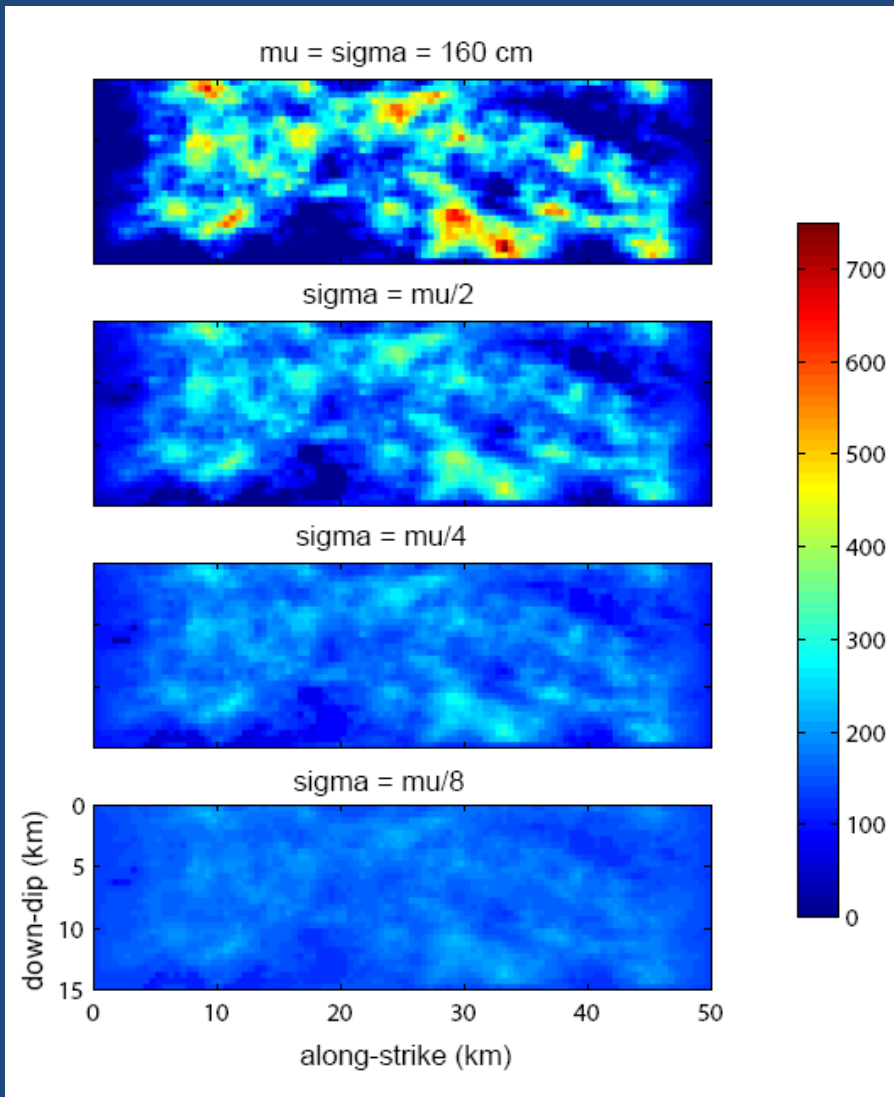


3) $\mathbf{X} \sim \mathcal{N}(\boldsymbol{\mu}, \boldsymbol{\Sigma})$

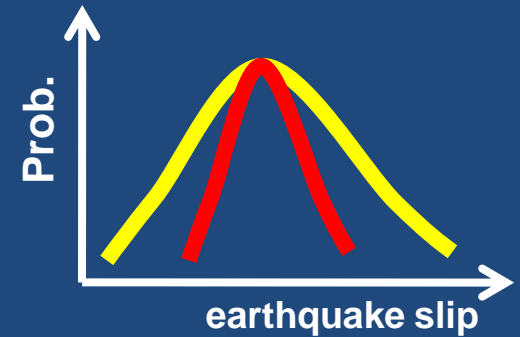
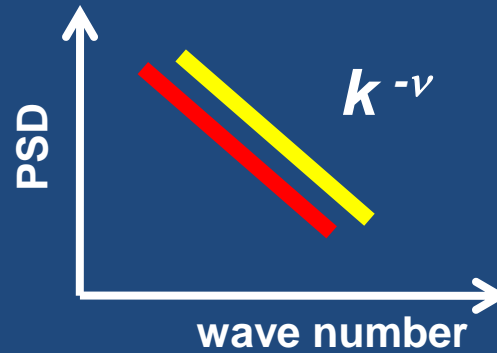
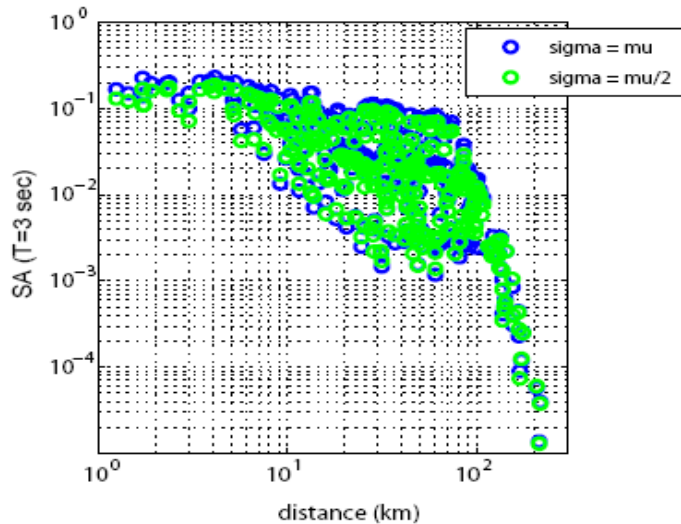
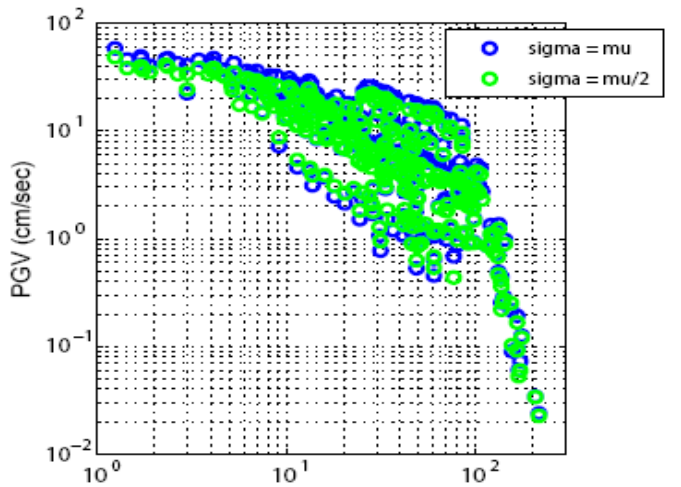


Slip Realizations

with the Same Spectral Decay (2-point stats)



Variability in Ground Motion with different 1-point statistics



Scaling law for sigma?

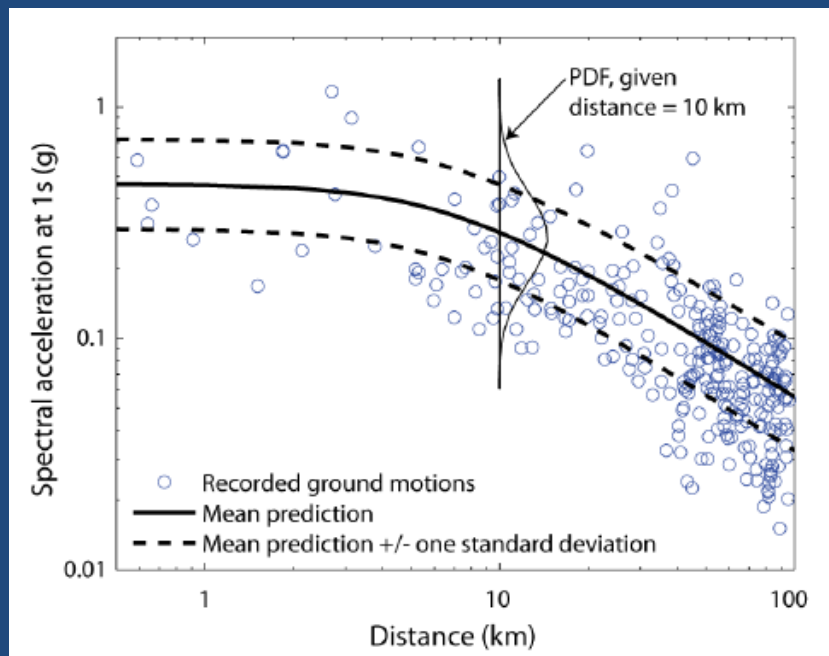
$\text{Log}(\mu) = a M + b$ [Somerville et al., 1999]

$\text{Log}(\sigma) = a M + b$ **missing!!**

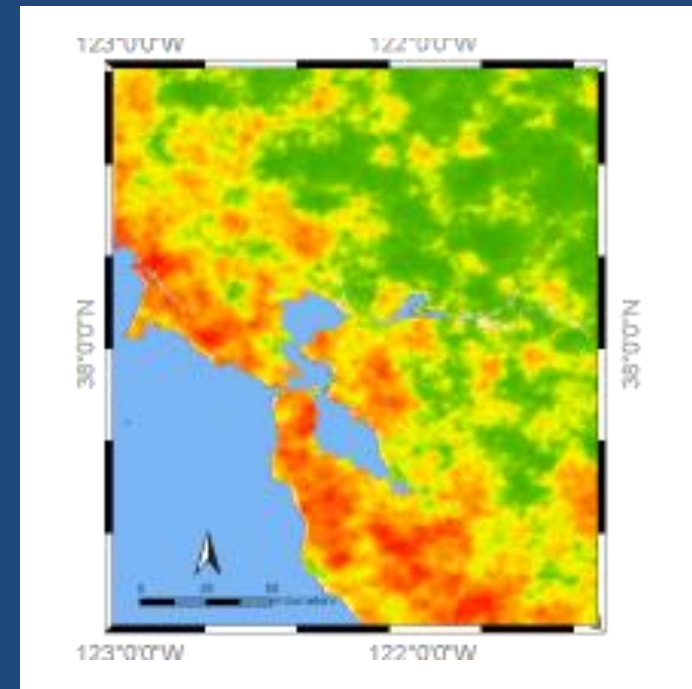
$\text{Log}(a_x) = a M + b$ [Mai and Beroza, 2002]

1-point and 2-point statistics *in Ground Motion Prediction*

1-point statistics in GMP

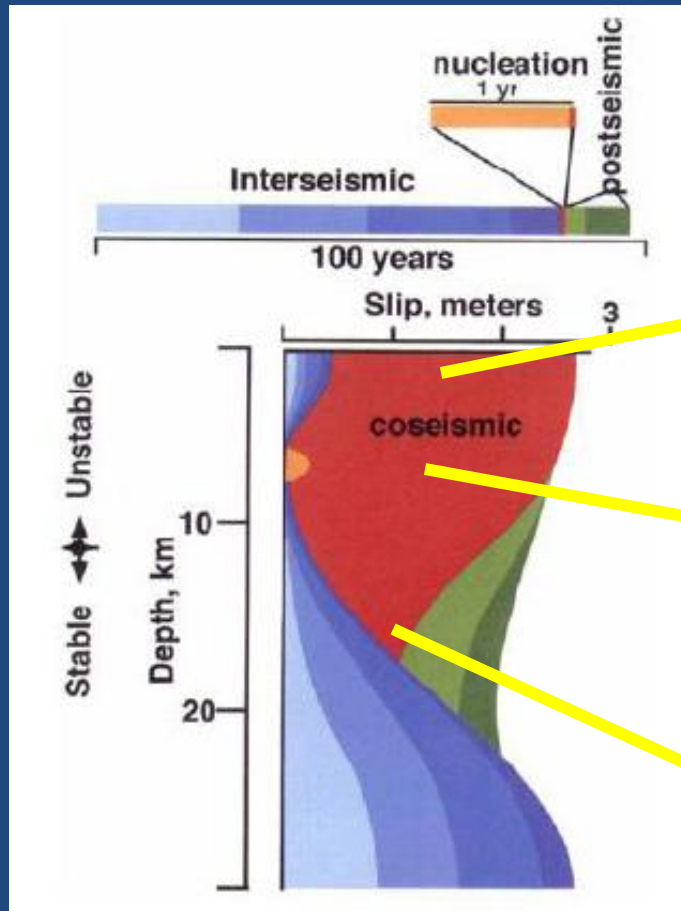


2-point statistics in GMP

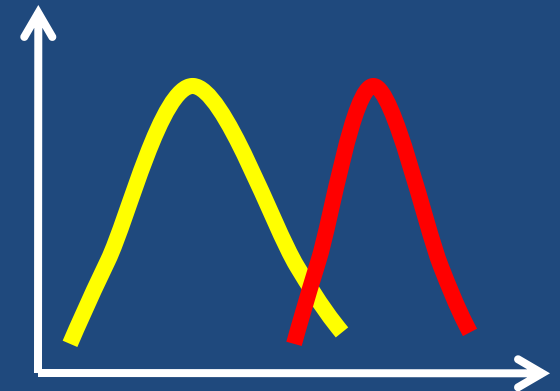


(Image source: J. Baker's website at Stanford Univ.)

Depth-dependency (Non-stationarity) of earthquake source statistics

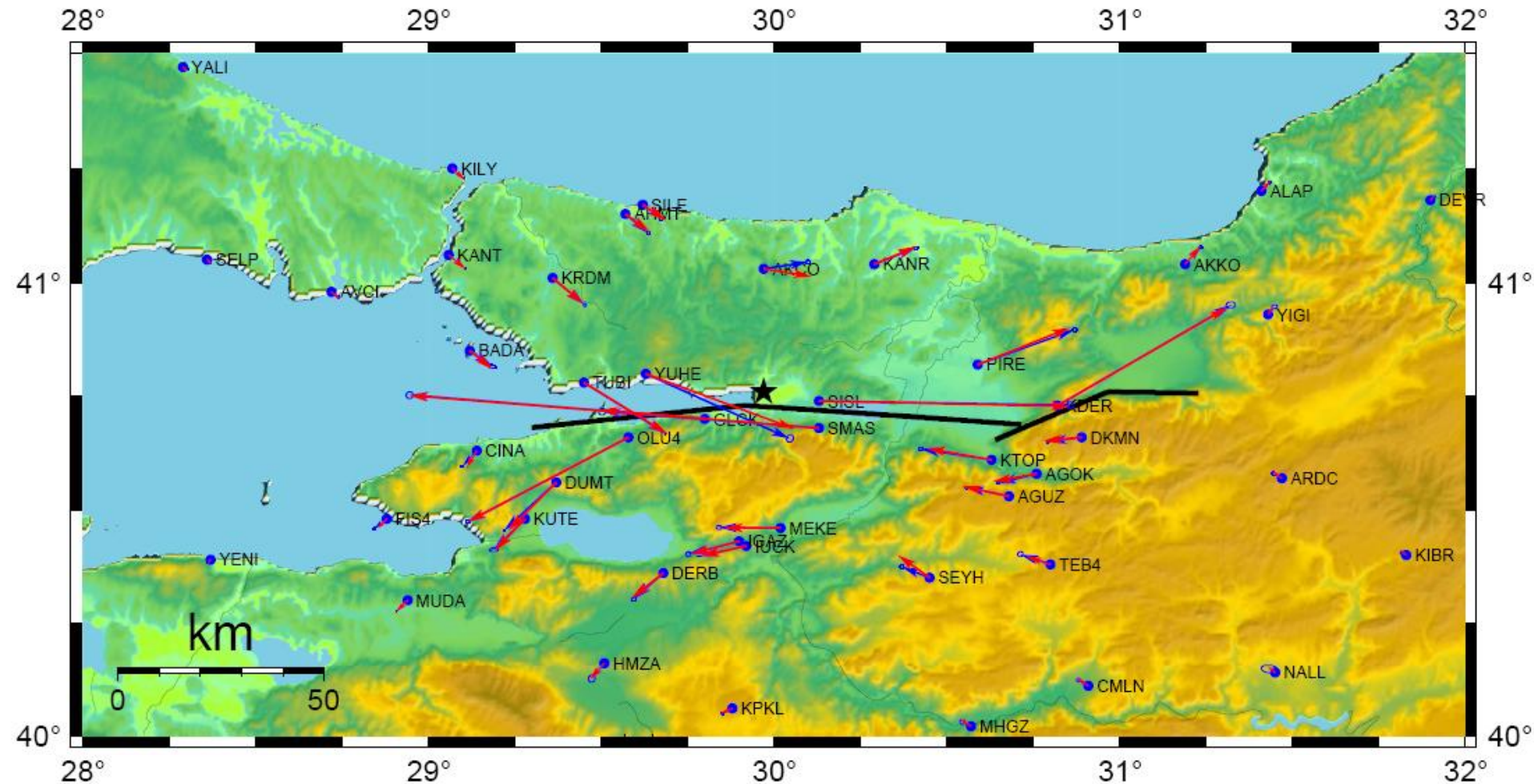


Slip Velocity Function



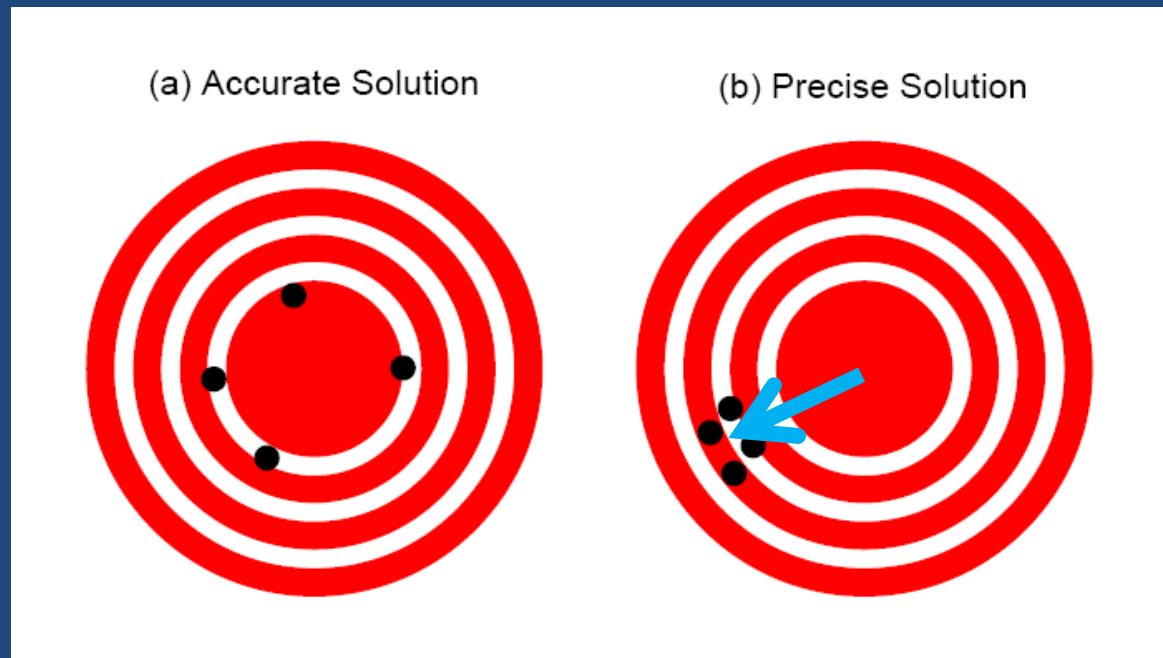
[Scholz, 2002]

Constraining 1-point and 2-point statistics *with Bayesian inversion*



(1999 Izmit, Turkey, event)

Accurate vs. Precise Solutions



$$\min \|\mathbf{d} - \mathbf{G}\mathbf{m}\|_2^2 + \alpha^2 \|\mathbf{L}\mathbf{m}\|_2^2, \hat{\mathbf{m}} = (\mathbf{G}^T \mathbf{G} + \alpha^2 \mathbf{L}^T \mathbf{L})^{-1} \mathbf{G}^T \mathbf{d}$$

$$E(\hat{\mathbf{m}}) \neq \mathbf{m}^{true} \Rightarrow \text{biased estimator!!}$$

Mean squared error:

$$MSE(\hat{\mathbf{m}}) = E((\hat{\mathbf{m}} - \mathbf{m}^{true})^2) = Var(\hat{\mathbf{m}}) + (Bias(\hat{\mathbf{m}}))^2$$

Tikhonov Regularization

- Advantage: improve the stability of inversion, otherwise very ill-posed inverse problems

$$\min \|\mathbf{d} - \mathbf{G}\mathbf{m}\|_2^2 + \alpha^2 \|\mathbf{L}\mathbf{m}\|_2^2, \hat{\mathbf{m}} = (\mathbf{G}^T\mathbf{G} + \alpha^2\mathbf{L}^T\mathbf{L})^{-1}\mathbf{G}^T\mathbf{d}$$

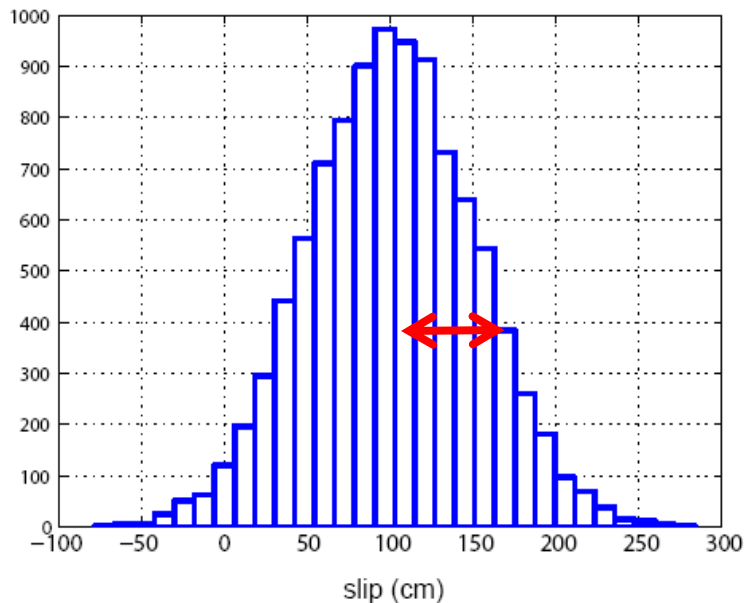
- By-products:
 - Lower resolution => inaccurate estimation of solutions
 - Biased => inaccurate estimation of uncertainty
 - Contaminates 1-point and 2-point stats of earthquake slip (and stress drop)

Prior model distribution

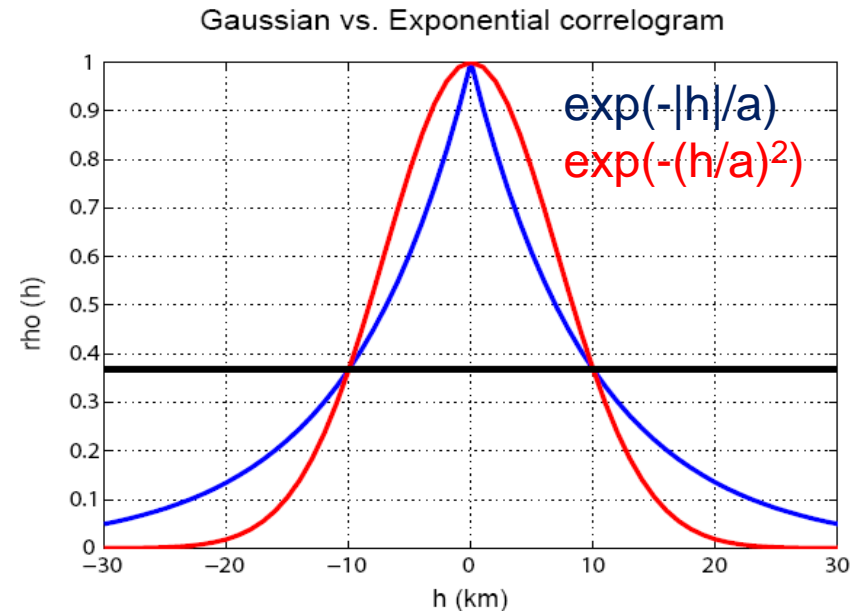
$$\sigma_M(\mathbf{m}) = k\rho_M(\mathbf{m}) \cdot \rho_D(\mathbf{g}(\mathbf{m}))$$

$$\rho_M(\mathbf{m}) = N(\boldsymbol{\mu}, \boldsymbol{\Sigma})$$

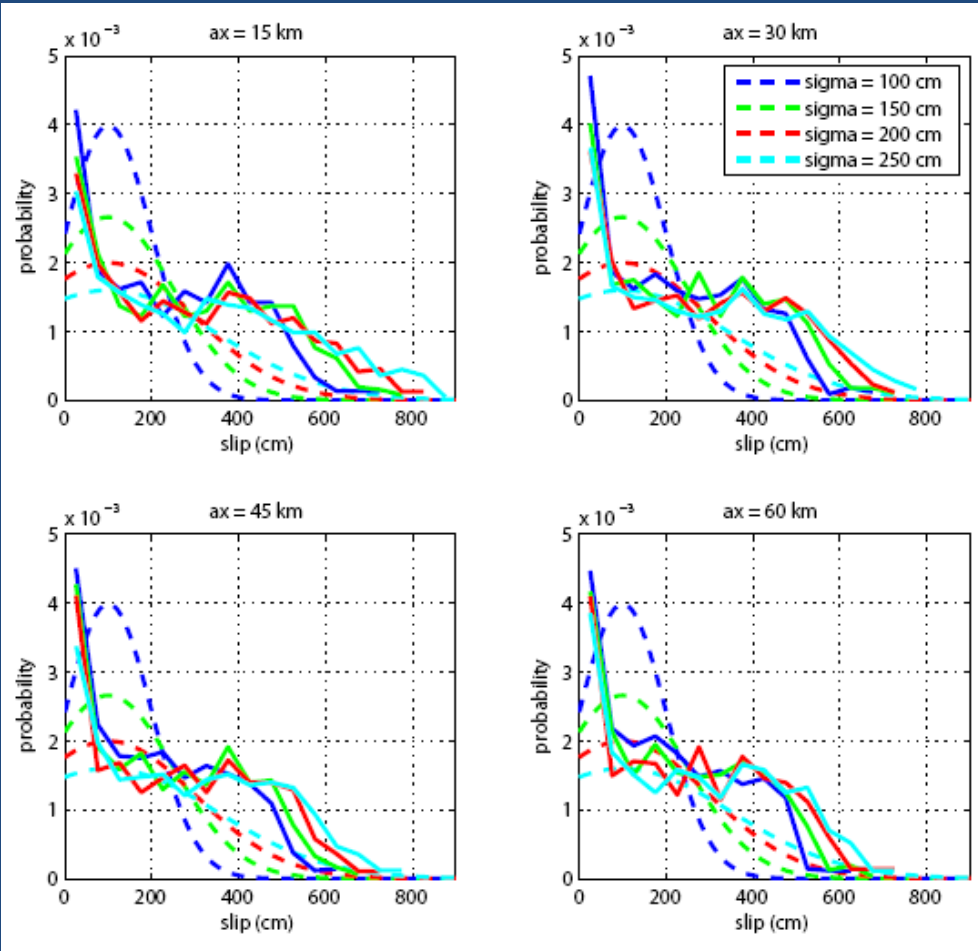
1-point statistics



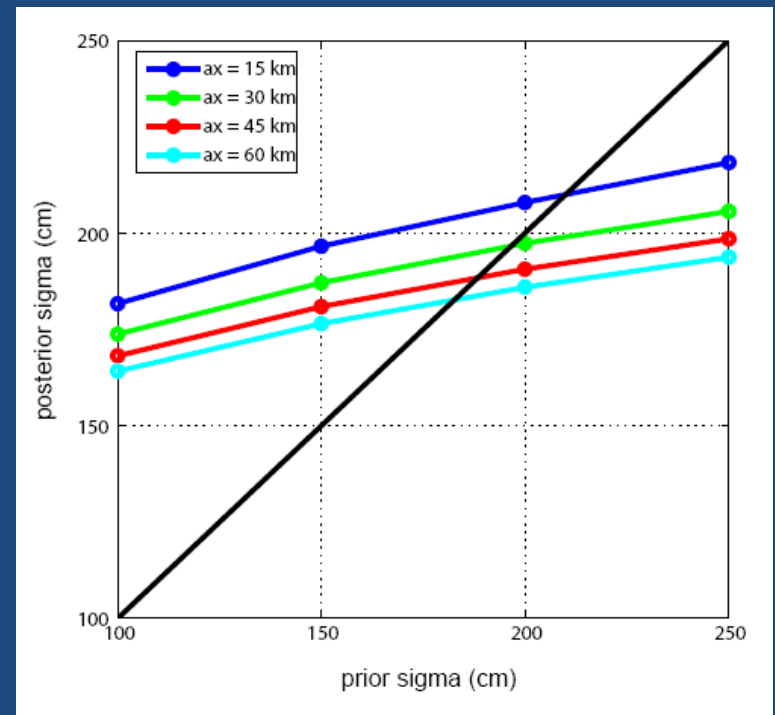
2-point statistics



1-Point Statistics



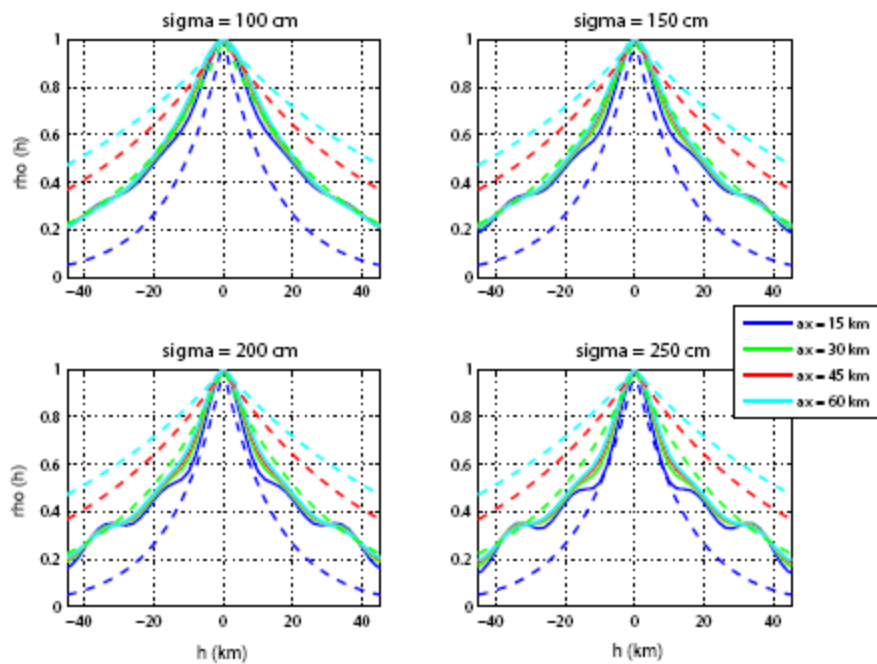
Dashed: prior
Solid: posterior



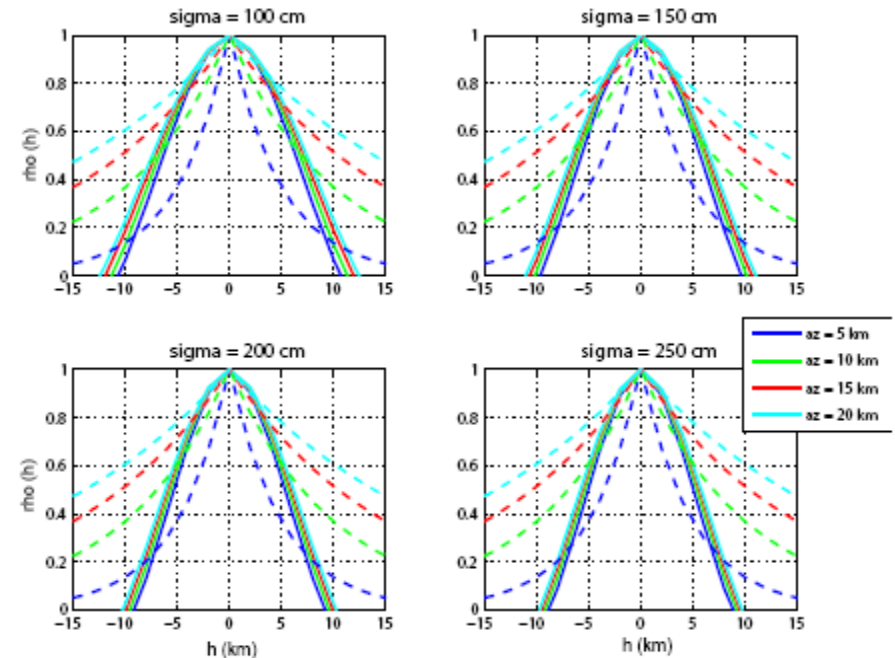
Spatial Coherence (2-point stats)

(dashed: prior, solid: posterior)

(a) along-strike

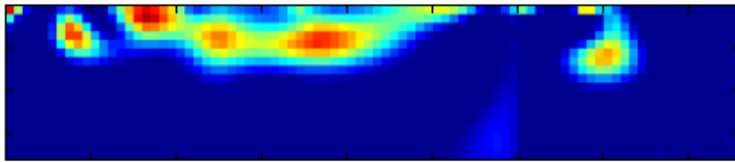


(b) down-dip

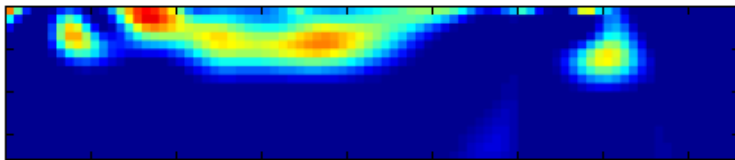


Estimated Slip Distributions

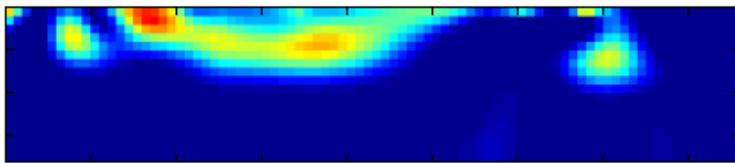
sigma = 200 cm, ax = 15 km



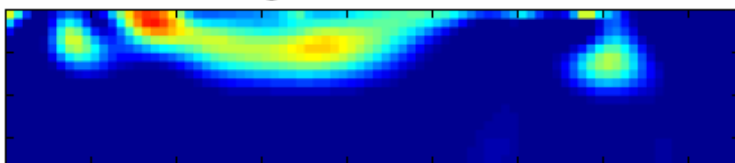
sigma = 200 cm, ax = 30 km



sigma = 200 cm, ax = 45 km

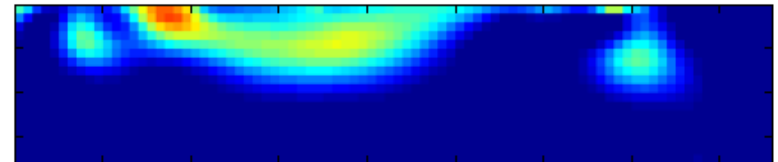


sigma = 200 cm, ax = 60 km

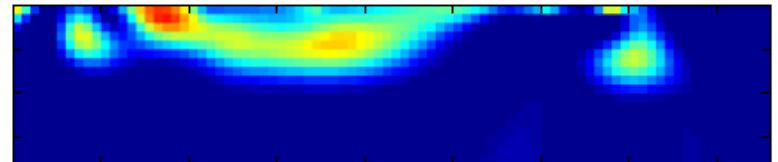


along-strike (km)

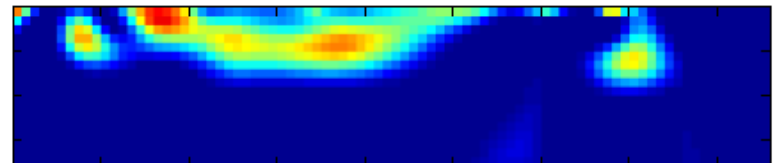
sigma = 100 cm, ax = 30 km



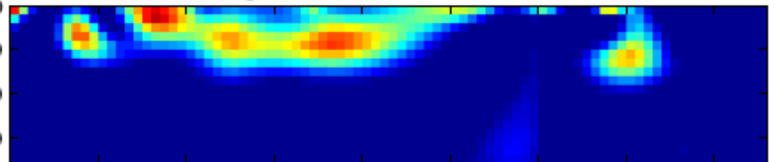
sigma = 150 cm, ax = 30 km



sigma = 200 cm, ax = 30 km



sigma = 250 cm, ax = 30 km



along-strike (km)

Inferring dynamic parameters *from kinematic rupture models*

$\Sigma =$

slip vs. slip	slip vs. V_r	slip vs. V_{max}	slip vs. rT
V_r vs. slip	V_r vs. V_r	V_r vs. V_{max}	V_r vs. rT
V_{max} vs. slip	V_{max} vs. V_r	V_{max} vs. V_{max}	V_{max} vs. rT
rT vs. slip	rT vs. V_r	rT vs. V_{max}	rT vs. rT

(Song and Somerville, 2010)

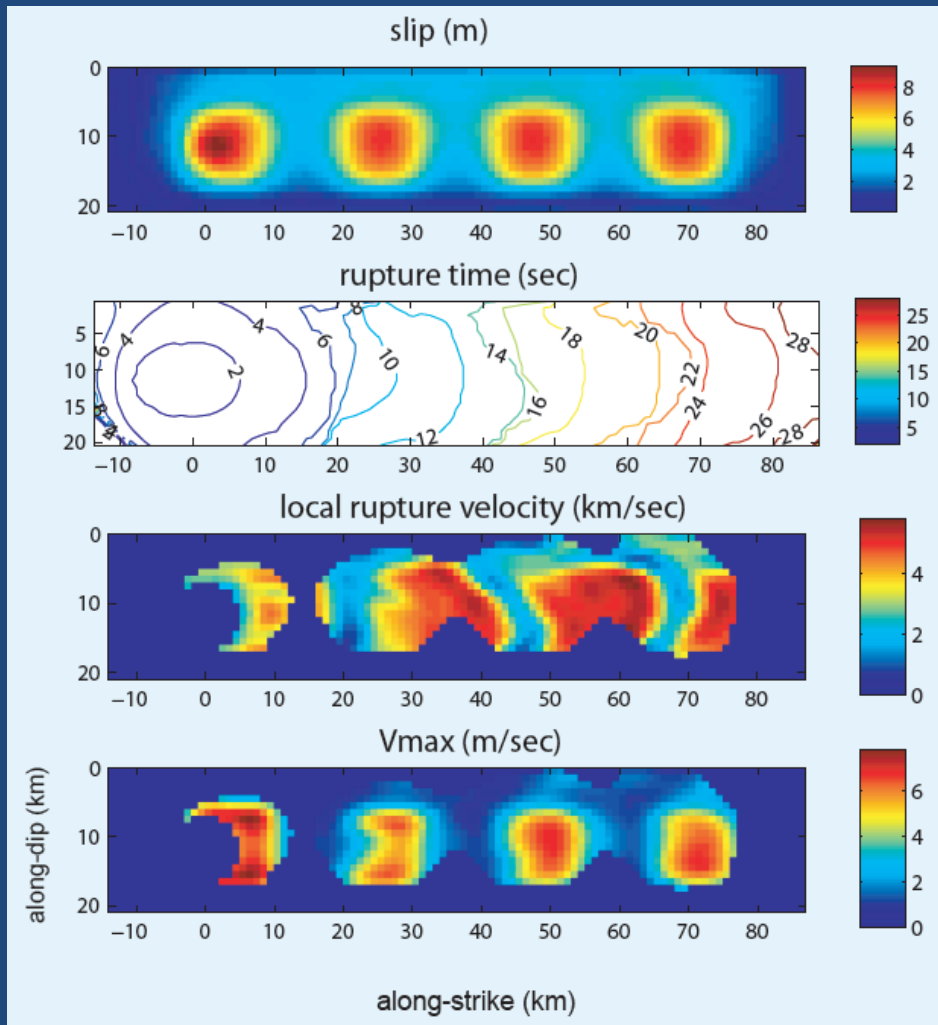
$\mathbf{u}(\underline{\mathbf{x}}, t) \Rightarrow \mathbf{T}(\underline{\mathbf{x}}, t) \Rightarrow sd, SE, d_c$

(Ide and Takeo, 1997; Tinti et al., 2005)

$\Sigma =$

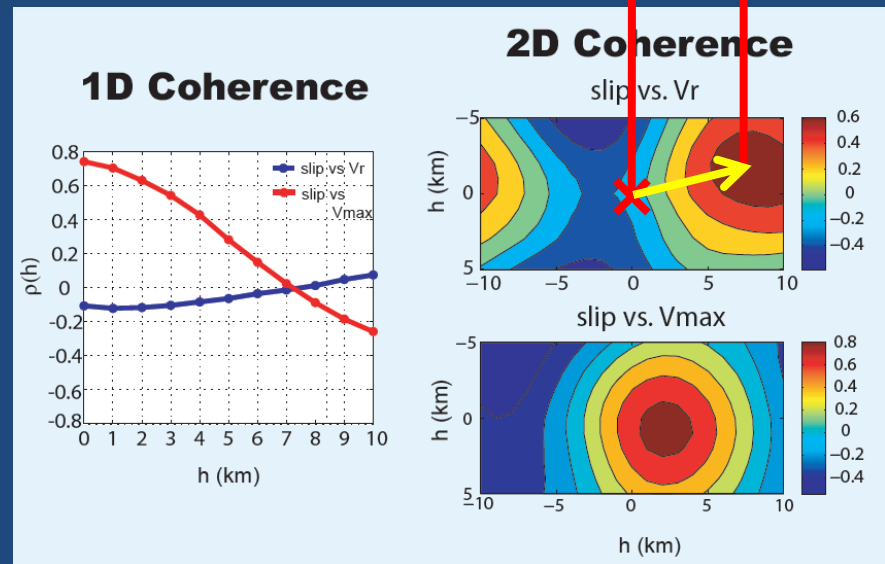
sd vs. sd	sd vs. G_c
G_c vs. sd	G_c vs. G_c

Spatial coherence from dynamic rupture models



(Dalguer et al., 2008)

Response Distance ~ 8 km



(Song and Somerville, 2010)

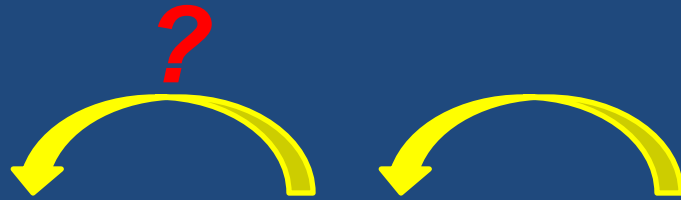
Summary (Part I)

- We should pay more attention to 1-point statistics and its non-stationarity.
- We may better constrain 1-point and 2-point statistics of source parameters by regularizing the model space with the same form of 1-point and 2-point statistics with Bayesian inversion.

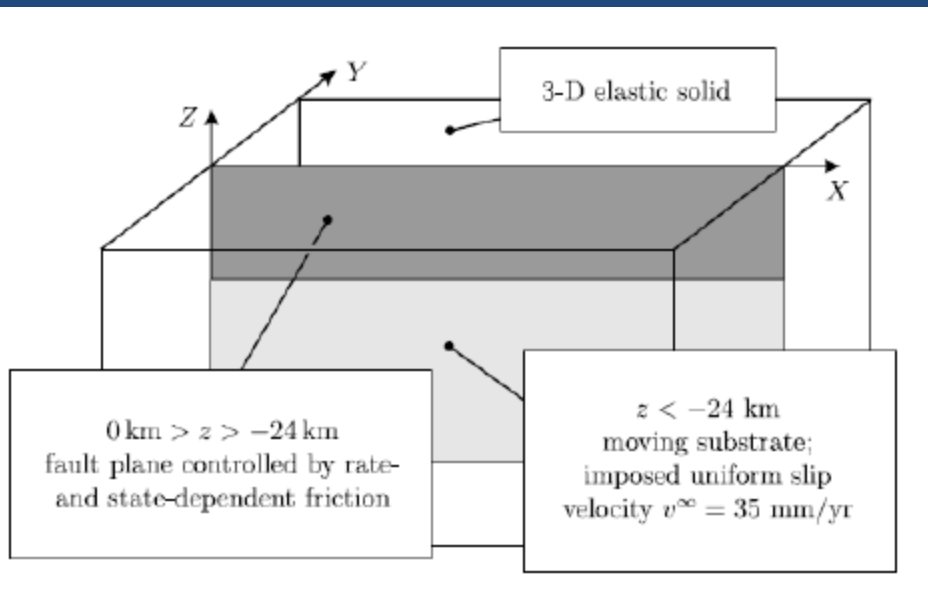
Outline

- **Spatial heterogeneity with 1-point and 2-point statistics**
 - Do we pay enough attention to 1-point statistics?
 - How to constrain them from data?
- **Two-step approach**
 - **Quasi-dynamic multi-cycle simulation with RS friction law**
 - **Full-dynamic single-event simulation with SW friction law**
 - Contributors: G. Hillers, A. Pitarka, P.M. Mai, L.A. Dalguer, P. Somerville
 - Supported by Japan Nuclear Energy Safety (JNES) through Geo Research Institute (GRI)

Quasi-dynamic multi-cycle simulation

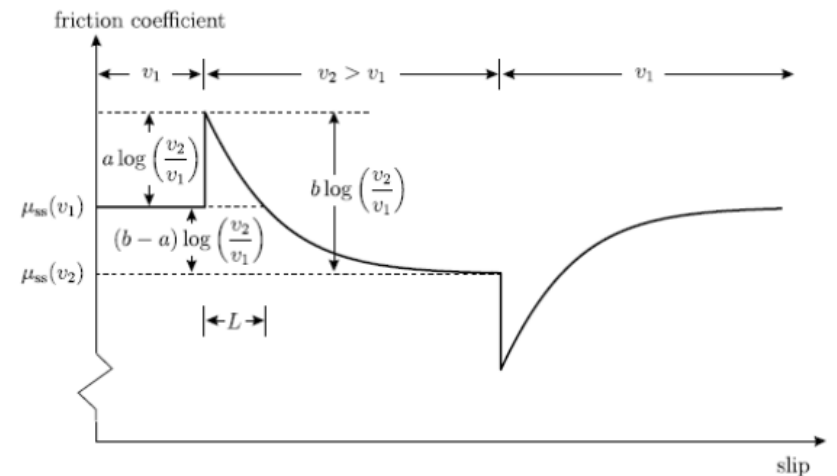


$a, b, L \Rightarrow sd, SE, d_c \Rightarrow slip, V_r, V_{max} \Rightarrow PGV, PGA, SA$



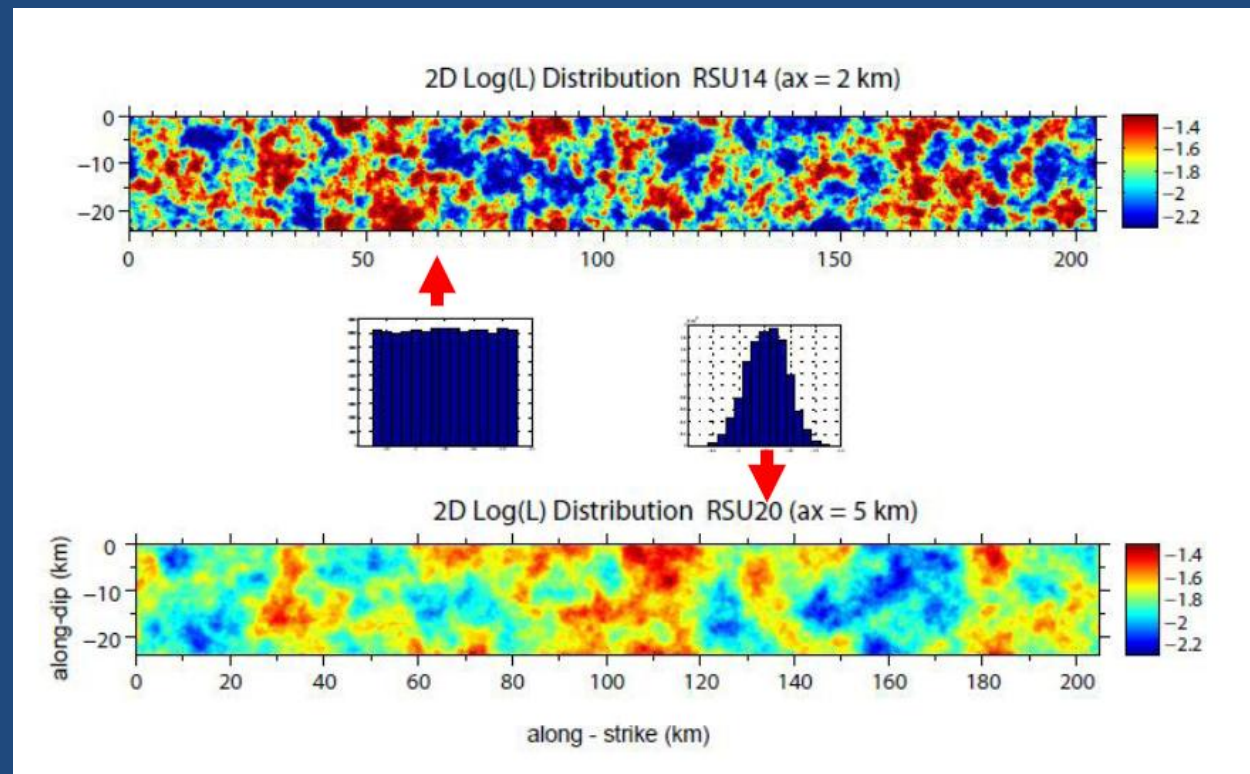
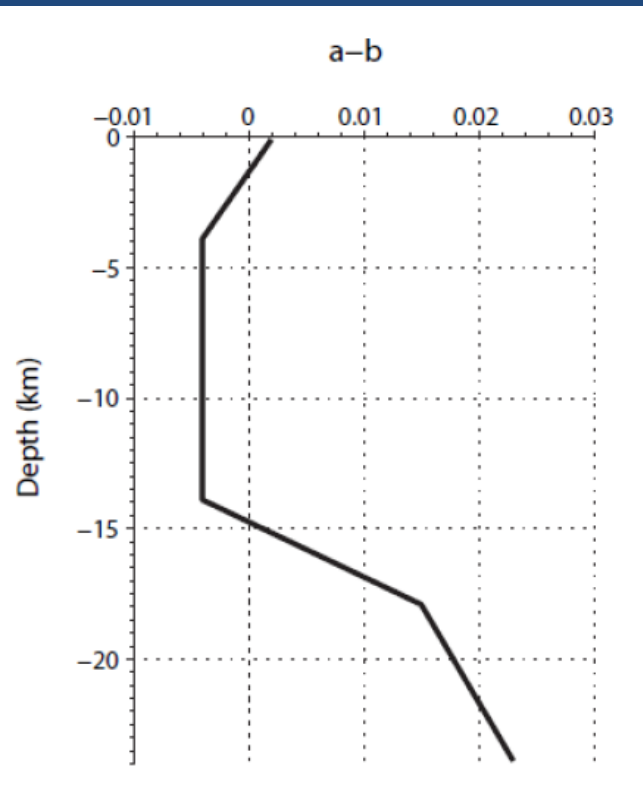
(model setup from Hillers et al., 2006)

Rate and state-dependent friction law

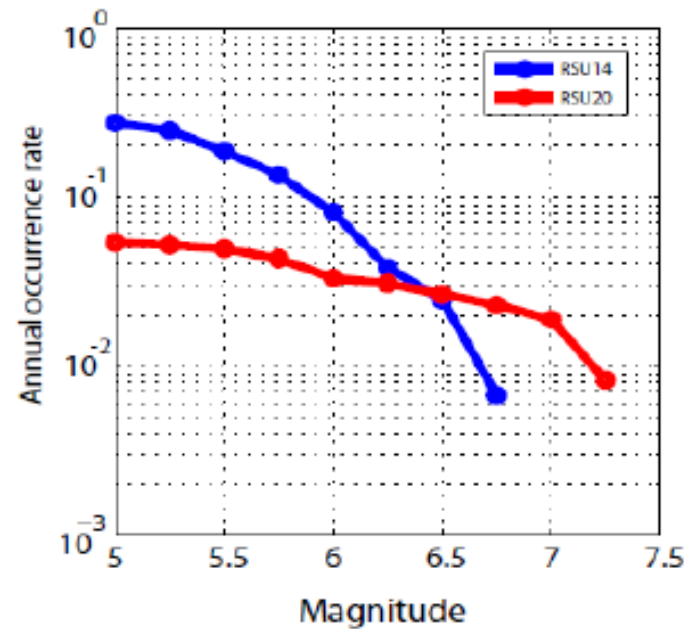
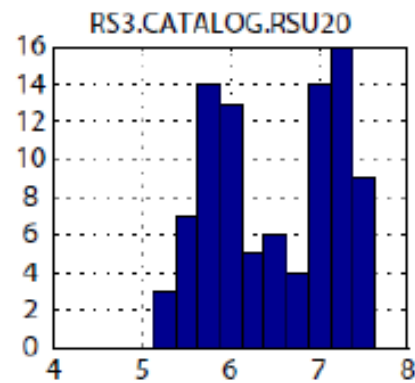
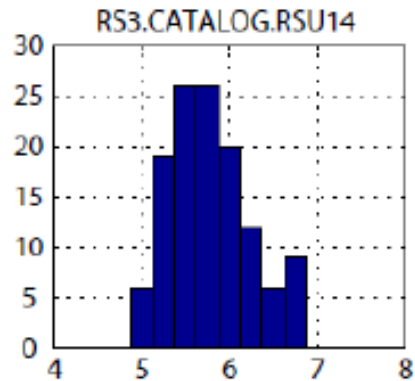


(Dieterich, 1979; Ruina, 1983)

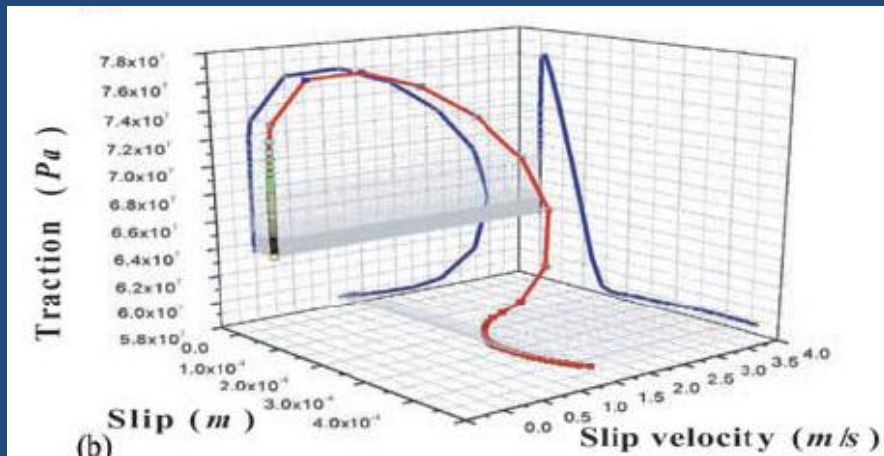
Input parameters (a , b , L) *in the multi-cycle simulation*



Gutenberg-Richter vs. Characteristics?

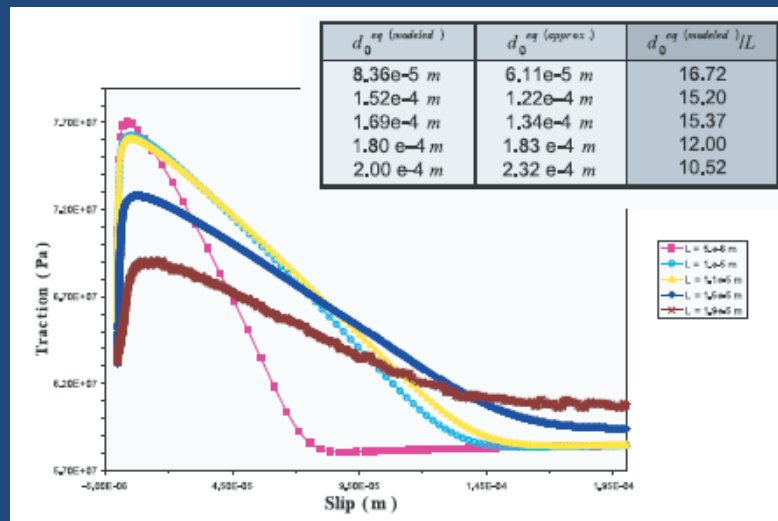


L to d_c conversion



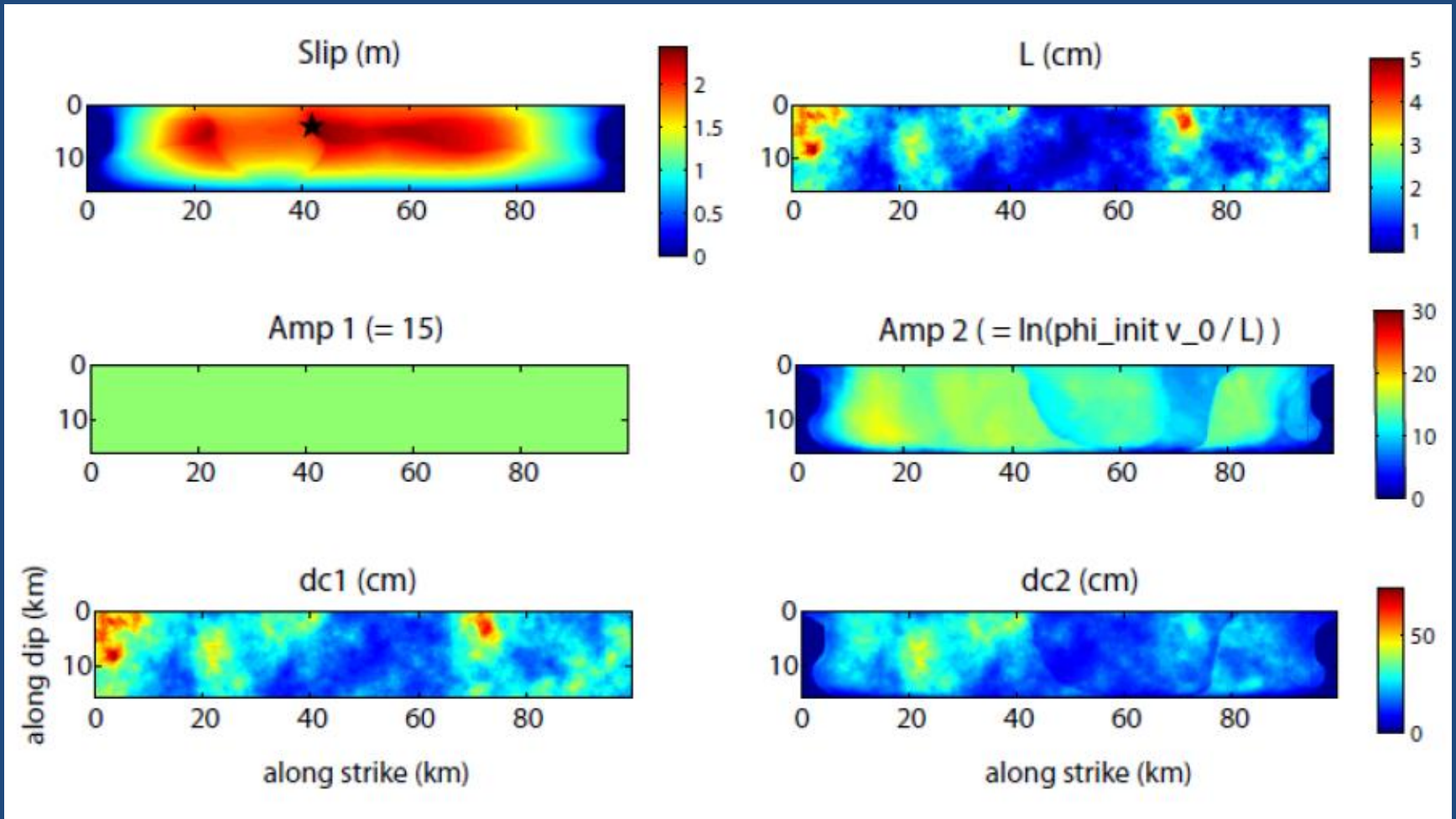
$$d_{c1} \cong L * 15$$

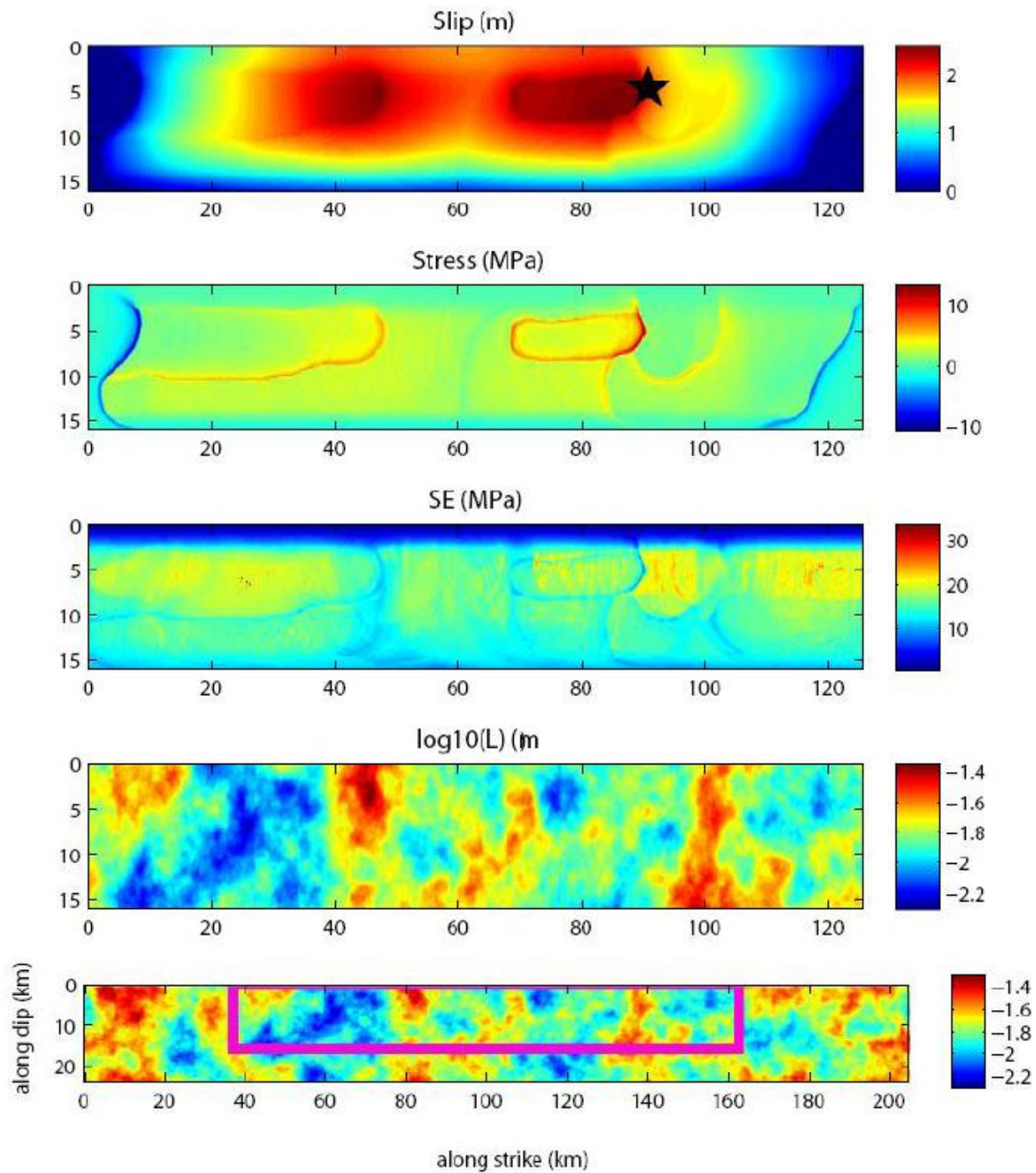
$$d_{c2} \cong L * \ln\left(\frac{\theta_{init} v_0}{L}\right)$$

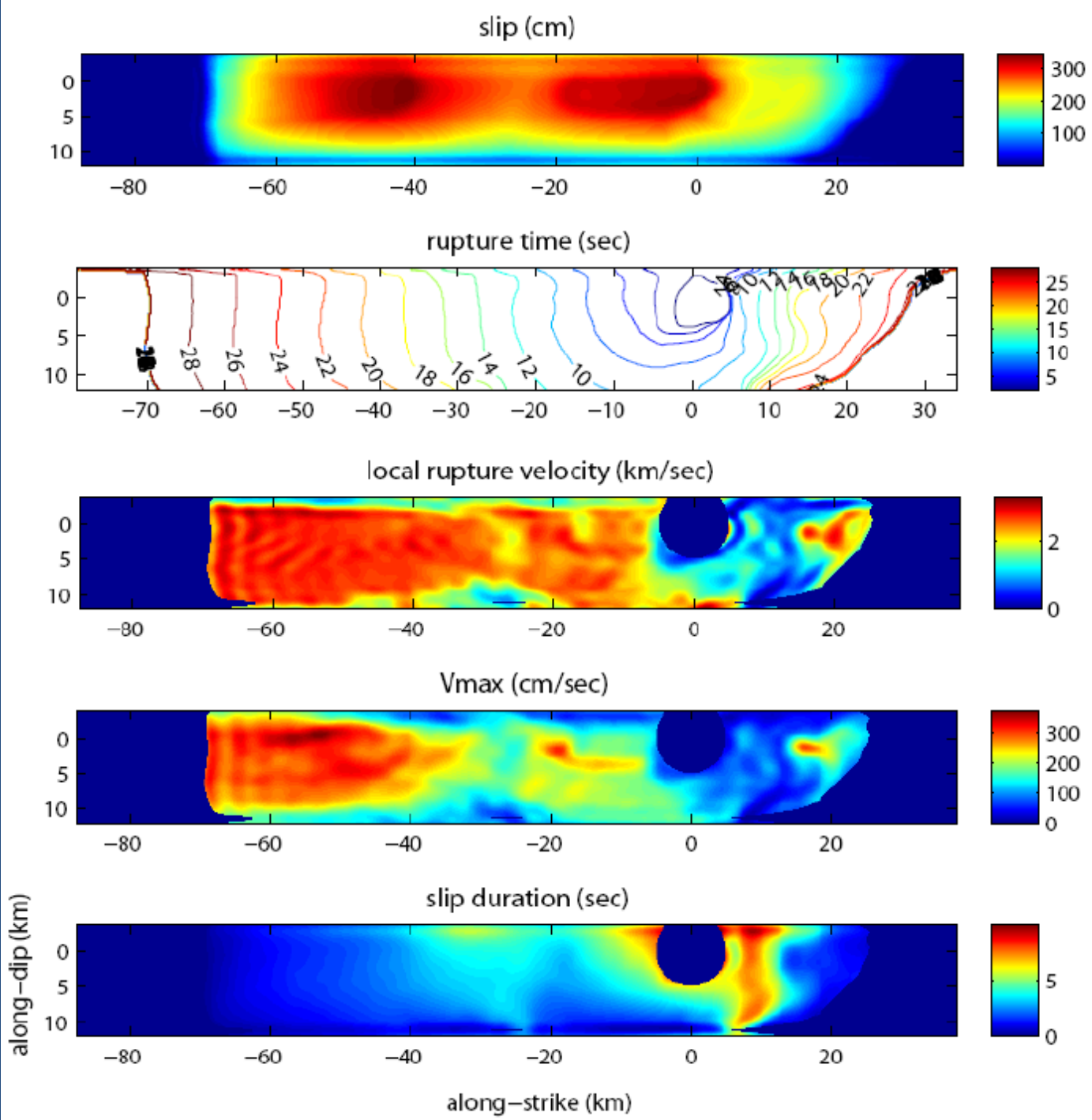


(Cocco and Bizzarri, 2002; Bizzarri and Cocco, 2003)

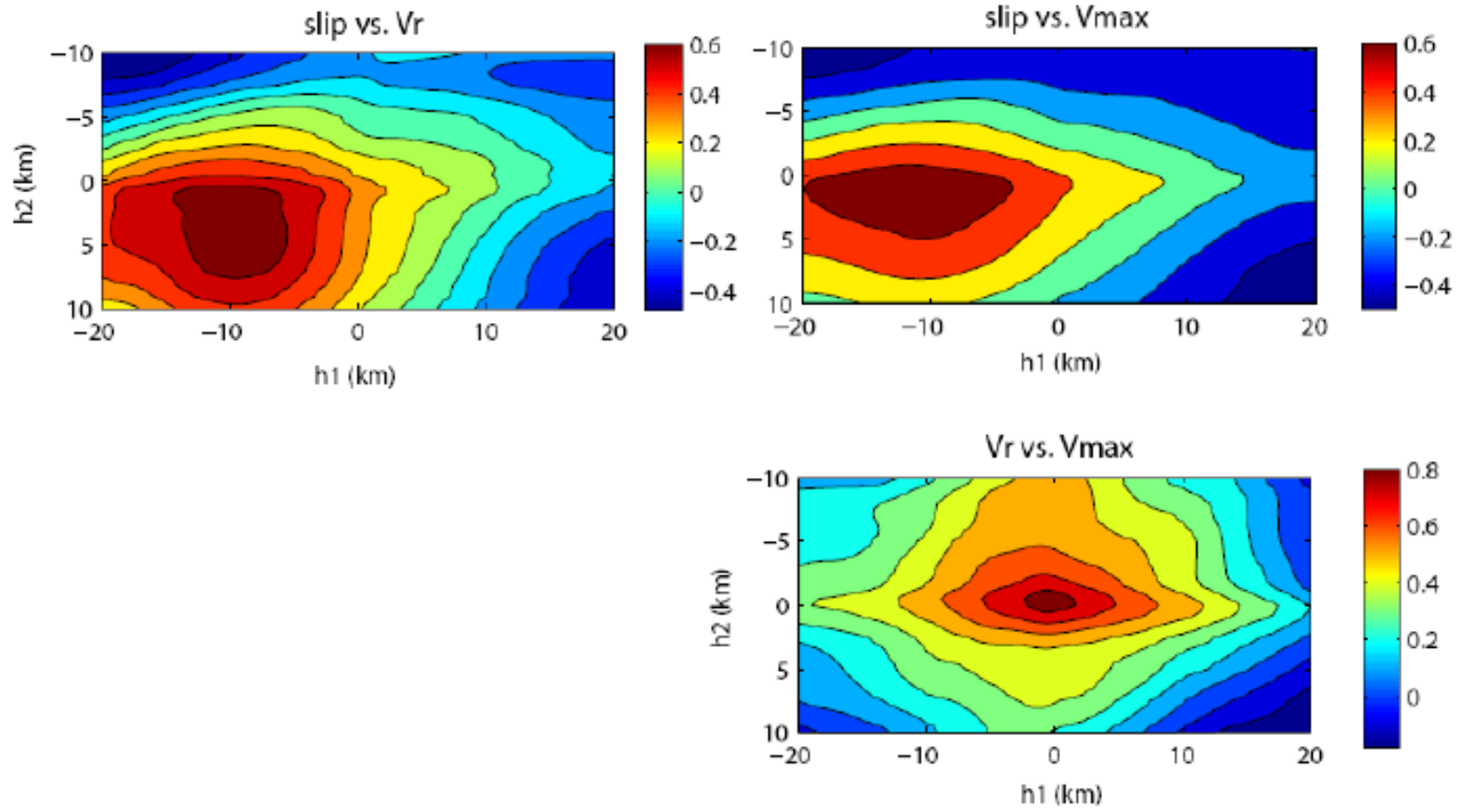
L to d_c conversion







Spatial cross-coherence



Summary (Part II)

- Physically self-consistent dynamic input parameters inferred from multi-cycle simulation
- Generate a series of events occurring on a single fault system through a cycle of the fault evolution
- **Applicability to “100-runs”**
 - Currently applied to strike-slip
 - Magnitude range