Inferences from the Field -Effects of Geometry on Earthquake Ruptures

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Surface Rupture Data

- Steps in surface rupture developed a large data set: 46 strike-slip, 30 dip-slip surface rupturing events.
- Bends in surface rupture: 67 rupture maps
- Annotated surface rupture maps available at BSSA.



Step and Gap Measurements



RESEARCH ARTICLE | MAY 24, 2016 Steps and Gaps in Ground Ruptures: Empirical Bounds on Rupture Propagation ©

Glenn P. Biasi; Steven G. Wesnousky Bulletin of the Seismological Society of America (2016) 106 (3): 1110-1124. https://doi.org/10.1785/0120150175

RESEARCH ARTICLE | OCTOBER 10, 2017 Bends and Ends of Surface Ruptures 🔗

Glenn P. Biasi; Steven G. Wesnousky Bulletin of the Seismological Society of America (2017) 107 (6): 2543-2560. https://doi.org/10.1785/0120160292

Step measurements only on surface fault sections 5-7 km minimum length.





Step sizes passed, strike-slip

Step size passed depends on rupture length.

- <5 km rule is too simple; big steps require big ruptures
- Shortest rupture passing 3 km step = 50 km long
- Shortest passing 4 km or more = 102 km long
- Interplay of dynamic/momentum effects vs. mechanical strength of step.

Event numbers from Biasi and Wesnousky, 2016

Strike-slip * Normal 0 15 69 40 Reverse ∇ 41 5 28 42 014 27 36 10² 10^{3}

Step sizes passed vs. rupture mechanism

- Dip slip events pass larger steps than strike slip. Up to 12 km.
- Surface rupture patterns show dip-slip ruptures respond to volumetric stresses

Surface Rupture Length

12

10

8

6

2

0

10

Step Size (km)

Event numbers from Biasi and Wesnousky, 2016



Passing Ratio

Passing ratio: for a given step size, what fraction of ruptures stop vs. what fraction continue through?

Observations:

- ruptures rarely stop at small steps: high passing ratio
- ruptures commonly stop at at large steps: low passing ratio
- Modeled as linear: PR=1.89-0.31*W for W=step width (km). Coin-toss at 3 km.
- Ratio should diverge up as step size -> 0.



Do Bends Arrest Surface Rupture?

- Summarize surface rupture in >=5-7 km linear reaches.
- Measure bends in surface rupture.
- Measure bends at ends where fault continues but rupture did not (these bends stopped rupture).



Maximum bend angles overcome in rupture

Strike-slip maximum internal bend ~30 degrees; exceptions have 3-D explanations

Dip-slip maximum interior bends form distinct population - half are >37 degrees

Upper: Maximum bend data for dip slip and strike slip Lower: Cumulative distribution by rupture mechanism



Ends of Ruptures

Measure bends of fault at rupture ends.

Dip slip rarely ends at a bend less than 20 degrees (clue to mechanics?)



Net Deflection: orientation difference between ending sections of the fault

Strike-slip: median difference is 11 degrees, and almost all are less than ~30 degrees.

Exceptions for strike-slip transition to dip slip (2 of 42 cases).

Most strike-slip ruptures consistent with a near-constant regional stress direction.

TAAD: Total absolute angle deviation



Total deflections and average curvature

Upper Figures: Sum absolute values of trend changes - how much the rupture changes direction.

Lower Figures: Average curvature,

- SS max = median 0.7 deg/km, max: 1.8 deg/km.
- Dip median 1.6 deg/km, max: 5.5 deg/km

Average curvature should scale with energy loss in dynamic rupture.

passing ratio for bend angles in ruptures

Strike slip

Passing Ratio

Dip slip



Strike slip:

- Small bend angles are crossed more than stop rupture
- Large bends stop ruptures more often ٠ than are passed. Because of the small data set, average

across different bin sizes

Summary:

- bends of 11 deg. are passed 2x • compared to stopping;
- 31 degrees stop 2x compared to ۲ passing.
- Strange pattern for dip slip. ۲

Probabilities of stopping or passing, strike-slip



Simple model of a fault bend



Empirical data suggest that excess friction at ~28 degrees roughly matches forces of dynamic rupture.

Conclusions

- Rich set of surface ruptures
 - Strike-slip
 - Maximum step jump depends on rupture length
 - Interior bends and net orientation change rarely exceed 30 degrees
 - Ruptures up to 10's of km can be explained by a constant stress orientation
 - Normal and Reverse
 - Dip-slip faults can rupture together in strange patterns if they share a common stress field.
 - Dip-slip fault step jumps often exceed 5 km.
 - Empirical data suggest excess resistance in bends of ~28 degrees balances dynamic rupture forces.
- Dynamic Modeling
 - Test/validate passing ratio conditions and probabilities.
 - Test rupture length dependence of maximum jump size.
 - Energetics of individual and average orientation changes.



2016 Kaikoura, New Zeland, M 7.8



Kaikoura, New Zealand, 2016

- How to anticipate this in seismic source characterization?
- Hope fault thought to be "ready", but did not break – misaligned with stresses on Kekerengu fault



Oblique moment tensor -> 3D effects.

Figure courtesy of Nicola Litchfield



1911 Chon-Kemin, Kyrgyzstan, M7.8

155 to 195 km of co-seismic rupture

25-40 km of overlapping rupture of opposite vergence to main rupture

Reverse rupture steps 10+ km across a mountain range in order to continue.

Circles: locations of slip estimates by Arrowsmith et al on 1911 rupture. View as field confirmation of 1911 rupture map

Arrowsmith et al., 2016



1911 Chon-Kemin, Kyrgyzstan

3 January 1911 **Chon-Kemin Earthquake Reverse Mechanism**



Reverse rupture across a mountain range in order to continue.

Vergence and rupture make sense in context of a single regional stress field.

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Fault continues west beyond west end of rupture, observed in displacement of Holocene sediments.

gap tabulation because of potential detection and preservation problems in alpine reach. Fault system continuity recognized in Delvaux et al 2001.

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Complex rupture: 1957 Gobi-Altai

1957 Gobi-Altai, Mongolia rupture

Nominally 235 km long strike slip

100's of km of reverse faulting with multi-meter offsets overlapping with the main trace.

^{onal} Simple triggering from one fault reach to ^{onal} the next is improbably.

Co-seismic rupture can be understood in the context of a regional stress field